

Did the COVID-19 Crisis Reduce Free Tropospheric Ozone across the Northern Hemisphere?

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55 **Key Points:**

56 • From April through August 2020, ozone stations in the northern extratropics report on
57 average 7% (or 4 ppbv) less ozone in the free troposphere than normal.

58 • Such low tropospheric ozone, over several months, and at so many sites, has not been
59 observed in any previous year since at least the year 2000.

60 • We suggest that most of the low tropospheric ozone in 2020 is a consequence of the
61 substantial emission reductions caused by decreased worldwide activity due to the
62 COVID-19 pandemic.

63

64 **Abstract**

65 Throughout spring and summer 2020, ozone stations in the northern extratropics recorded
66 unusually low ozone in the free troposphere. From April to August, and from 1 to 8 kilometers
67 altitude, ozone was on average 7% (≈ 4 ppbv) below the 2000 to 2020 climatological mean. Such
68 low ozone, over several months, and at so many stations, has not been observed in any previous
69 year since at least 2000. Atmospheric composition re-analyses from the Copernicus Atmosphere
70 Monitoring Service and simulations from the NASA GMI model indicate that the large 2020
71 springtime ozone depletion in the Arctic stratosphere has contributed less than one quarter to the
72 observed tropospheric anomaly. The observed anomaly is consistent with two recent model
73 simulations, which assume emission reductions similar to those caused by the COVID-19 crisis.
74 COVID-19 related emission reductions appear to be the major cause for the observed low free
75 tropospheric ozone in 2020.

76

77 **Plain Language Summary**

78 Worldwide actions to curb the spread of the COVID-19 virus have closed factories, grounded
79 airplanes, and have generally reduced travel and transportation. Less fuel was burnt, and less
80 exhaust was emitted into the atmosphere. Due to these measures, the concentration of nitrogen
81 oxides and volatile organic compounds (VOCs) decreased in the atmosphere. These substances
82 are important for photochemical production and destruction of ozone in the atmosphere. In clean
83 or mildly polluted air, reducing nitrogen oxides and/or VOCs will reduce the photochemical
84 production of ozone and result in less ozone. In heavily polluted air, in contrast, reducing
85 nitrogen oxides can increase ozone concentrations, because less nitrogen oxide is available to
86 destroy ozone. In this study, we use data from three types of ozone instruments, but mostly from
87 ozonesondes on weather balloons. The sondes fly from the ground up to 30 kilometers altitude.
88 In the first 10 kilometers we find significantly reduced ozone concentrations in spring and
89 summer of 2020, less than in any other year since at least 2000. We suggest that reduced
90 emissions due to the COVID-19 crisis have lowered photochemical ozone production and have
91 caused the observed ozone reductions in the first 10 kilometers of the atmosphere, the
92 troposphere.

93

94 **1 Introduction**

95 Widespread slowdowns caused by the COVID-19 pandemic have reduced anthropogenic
96 emissions throughout the year 2020. Guevara et al. (2020) report emission reductions up to 60%
97 for NO_x , and up to 15% for non-Methane Volatile Organic Compounds (NMVOC) over Europe
98 for March and April 2020 (Barré et al., 2020). Based on satellite observations of NO_2 columns
99 (Bouwens et al., 2020), comparable NO_x emission reductions are reported for Chinese cities
100 during February 2020 (Ding et al., 2020; Feng et al., 2020). For the first half of 2020, Liu et al.
101 (2020) report an overall reduction of 8.8% for CO_2 emissions, consistent in magnitude with the
102 mentioned NO_2 emission reductions. The largest relative reductions occurred for airtraffic, where
103 traffic (and emissions) decreased by $\approx 40\%$ in the first half of 2020 (Liu et al., 2020), and have
104 remained low during the second half of 2020.

105 COVID 19 emission reductions are large enough to affect ozone levels in the troposphere
106 (Dentener et al., 2011). Tropospheric O₃-NO_x-VOC-HO_x chemistry is, however, complex and
107 non-linear. The net effect of emission changes on ozone depends on NO_x and VOC
108 concentrations, and on their ratios (Kroll et al., 2020; Sillman, 1999; Thornton et al., 2002). In
109 polluted regions, at high NO_x concentrations (>> 1ppb), reducing NO_x concentrations can
110 increase ozone, because ozone titration by NO is reduced (Sicard et al., 2020). At low
111 concentrations (NO_x < 1ppb), however, in the clean or mildly polluted free troposphere, reducing
112 NO_x lowers photochemical ozone production (Bozem et al., 2017) and results in less ozone.

113 Indeed, for many polluted regions, studies report increased near-surface ozone
114 concentrations after COVID-19 lockdowns (Collivignarelli et al., 2020; Shi & Brasseur, 2020;
115 Siciliano et al., 2020; Venter et al., 2020). Reduced surface ozone is reported for some rural areas
116 after COVID-19 lockdowns, e.g., in the US and Western Europe (Chen et al., 2020; Menut et al.,
117 2020). Meteorological conditions complicate matters, and play an important role as well
118 (Goldberg et al., 2020; Keller et al., 2020; Ordóñez et al., 2020).

119 In this paper we report significant ozone reductions observed in the free troposphere at
120 many stations in the northern extratropics. These large-scale reductions occurred in late spring
121 and summer 2020, following the widespread COVID-19 slowdowns, and are unique for the last
122 two decades.

123 **2 Instruments and Data**

124 Regular observations of ozone in the free troposphere are sparse: Only around 50 ozone
125 sounding stations worldwide (e.g. Tarasick et al., 2019), a handful of tropospheric lidars (Gaudel
126 et al., 2015; Granados-Muñoz and Leblanc 2016; Leblanc et al., 2018), and about twenty Fourier
127 Transform Infrared Spectrometers (FTIRs, Vigouroux et al., 2015). In-Service Aircraft for a
128 Global Observing System (IAGOS, Nédélec et al., 2015) are another important source of
129 tropospheric ozone data. Due to the COVID-19 slowdowns, however, few IAGOS aircraft were
130 flying in 2020, and IAGOS data became quite sparse. The information content of satellite
131 measurements on ozone in the free troposphere is limited: Typically, only one value (one degree
132 of freedom) for the entire troposphere, with modest accuracy, 10 to 30% (Hurtmans et al., 2012;
133 Liu et al., 2010; Oetjen et al., 2014). The recent Tropospheric Ozone Assessment Report found
134 large differences in tropospheric ozone trends derived from different satellite instruments, and
135 even different signs in some regions (Gaudel et al., 2018).

136 Ozonesondes measure profiles with high vertical resolution, about 100 m, and good
137 accuracy, about 5 to 15% in the troposphere, 5% in the stratosphere (Smit et al., 2007; Sterling et
138 al., 2018). This is adequate to detect ozone anomalies of several percent. Substantial work has
139 gone into standardizing and improving operating procedures for ozonesondes (WMO, 2014).
140 Homogenization of historical records has started as well (Tarasick et al., 2016; Van Malderen et
141 al., 2016; Witte et al., 2017; Sterling et al., 2018). We use stations with regular soundings, at
142 least once per month since the year 2000, and with data available until at least July 2020.
143 Soundings with obvious deficiencies were rejected (large data gaps, ozone column from the
144 sounding deviating by more than 30% from ground- or satellite-based measurement). Table 1
145 provides information on stations, and public data archives.

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147

148 **Table 1.** Stations in this study, mostly ozonesonde stations. *FTIR and LIDAR stations are*
 149 *italicized.* Data sources: **W**=World Ozone and UV Data Centre
 150 (https://woudc.org/archive/Archive-NewFormat/OzoneSonde_1.0_1/), **N**=Network for the
 151 Detection of Atmospheric Composition Change (<ftp://ftp.cpc.ncep.noaa.gov/ndacc/station/>;
 152 <ftp://ftp.cpc.ncep.noaa.gov/ndacc/RD/>), **E**= European Space Agency Validation Data Center
 153 (<https://evdc.esa.int/> requires registration, or
 154 <ftp://zardozi.nilu.no/nadir/projects/vintersol/data/o3sondes> requires account), **G**=Global
 155 Monitoring Laboratory, National Oceanic and Atmospheric Administration
 156 (<ftp://aftp.cmdl.noaa.gov/data/ozwv/Ozonesonde/>)

157 ¹ Currently, Canadian data are available only up to March or April 2020. Newer Canadian data should become
 158 available for the final version of this study, and will be included. Newer data from other stations will be included as
 159 well for the final version.

160 ² Tateno data were corrected for the change from Carbon Iodine to ECC ozonesondes in December 2009.

161 ³ Stations affected by a drop-off in ECC sonde sensitivity > 3% in the stratosphere, after 2015 (see Stauffer et al.,
 162 2020). The drop-off is much smaller (<< 1%) in the troposphere, and should be negligible here. At many of the
 163 affected stations, ECC sondes behaved normally again in 2019/2020.

164

Station	Latitude (deg N)	Longitude (deg E)	Data source (see caption)	Data until	Profiles / spectra per month in 2020
Alert, Canada ^{1,3}	82.50	-62.34	W	4/2020	3.75
Eureka, Canada ^{1,3}	80.05	-86.42	W, E	4/2020	5.75
Ny-Ålesund, Norway	78.92	11.92	W, E	8/2020	7.63
<i>Ny-Ålesund FTIR, Norway</i>	78.92	11.92	<i>N</i>	7/2020	12.86
<i>Thule FTIR, Greenland</i>	76.53	-68.74	<i>N</i>	9/2020	73
Resolute, Canada ¹	74.72	-94.98	W	4/2020	5.50
Scoresbysund, Greenland	70.48	-21.95	E	9/2020	3.89
<i>Kiruna FTIR, Sweden</i>	67.41	20.41	<i>N</i>	7/2020	46
Sodankylä, Finland	67.36	26.63	W, E	8/2020	3.00
Lerwick, United Kingdom	60.13	-1.18	W, E	8/2020	4.38
Churchill, Canada ^{1,3}	58.74	-93.82	W	3/2020	3.33
Edmonton, Canada ^{1,3}	53.55	-114.10	W	3/2020	3.67
Goose Bay, Canada ¹	53.29	-60.39	W	3/2020	2.67
<i>Bremen FTIR, Germany</i>	53.13	8.85	<i>N</i>	10/2020	5.27
Legionowo, Poland	52.40	20.97	W	10/2020	4.00
Lindenberg, Germany	52.22	14.12	W	10/2020	4.60
DeBilt, Netherlands	52.10	5.18	W, E	8/2020	4.25
Valentia, Ireland	51.94	-10.25	W, E	8/2020	2.75
Uccle, Belgium	50.80	4.36	W, E	8/2020	12.13

Hohenpeissenberg, Germany	47.80	11.01	W	10/2020	10.10
<i>Zugspitze FTIR, Germany</i>	<i>47.42</i>	<i>10.98</i>	<i>N</i>	<i>9/2020</i>	<i>73</i>
<i>Jungfrauoch FTIR, Switzerland</i>	<i>46.55</i>	<i>7.98</i>	<i>N</i>	<i>10/2020</i>	<i>49</i>
Payerne, Switzerland	46.81	6.94	W	10/2020	11.10
Haute Provence, France	43.92	5.71	N	8/2020	2.50
<i>Haute Provence LIDAR, France</i>	<i>43.92</i>	<i>5.71</i>	<i>N</i>	<i>8/2020</i>	<i>3.50</i>
<i>Toronto FTIR, Canada</i>	<i>43.66</i>	<i>-79.40</i>	<i>N</i>	<i>10/2020</i>	<i>59</i>
Trinidad Head, California, USA	41.05	-124.15	G	8/2020	4.00
Madrid, Spain	40.45	-3.72	W	10/2020	4.10
Boulder, Colorado, USA	39.99	-105.26	G	8/2020	5.13
<i>Boulder FTIR, Colorado, USA</i>	<i>39.99</i>	<i>-105.26</i>	<i>N</i>	<i>10/2020</i>	<i>56</i>
Tateno (Tsukuba), Japan ²	36.05	140.13	W	6/2020	3.50
<i>Table Mountain LIDAR, California, USA</i>	<i>34.40</i>	<i>-117.70</i>	<i>N</i>	<i>8/2020</i>	<i>19</i>
Izana, Tenerife, Spain	28.41	-16.53	W	8/2020	2.00
<i>Izana FTIR, Tenerife, Spain</i>	<i>28.30</i>	<i>-16.48</i>	<i>N</i>	<i>9/2020</i>	<i>28</i>
Hong Kong, China	22.31	114.17	W	9/2020	4.11
Hilo, Hawaii, USA ³	19.72	-155.07	G	8/2020	4.00
<i>Mauna Loa FTIR, Hawaii, USA</i>	<i>19.54</i>	<i>-155.58</i>	<i>N</i>	<i>10/2020</i>	<i>36</i>
Paramaribo, Suriname	5.81	-55.21	N, E	9/2020	3.56
Pago Pago, American Samoa ³	-14.25	-170.56	G	9/2020	2.67
Suva, Fiji ³	-18.13	178.32	G	9/2020	1.44
<i>Wollongong FTIR, Australia</i>	<i>-34.41</i>	<i>150.88</i>	<i>N</i>	<i>10/2020</i>	<i>43</i>
Broadmeadows, Australia	-37.69	144.95	W	7/2020	4.29
Lauder, New Zealand	-45.04	169.68	W	10/2020	4.40
<i>Lauder FTIR, New Zealand</i>	<i>-45.04</i>	<i>169.68</i>	<i>N</i>	<i>10/2020</i>	<i>99</i>
Macquarie Island, Australia	-54.50	158.94	W	7/2020	4.29

165

166 Apart from the sondes, FTIR spectrometers from the Network for the Detection of
167 Atmospheric Composition Change (NDACC, De Mazière et al., 2018) provide independent
168 information, based on a completely different method (ground-based solar-infrared absorption
169 spectrometry). Altitude resolution of FTIR ozone profiles in the troposphere is much coarser (5
170 to 10 km) compared to the sondes, while accuracy is similar, 5 to 10% (Vigouroux et al., 2015).
171 Finally, we use data from tropospheric lidars (Gaudel et al., 2015, Granados-Muñoz & Leblanc
172 2016), which provide ozone profiles from ≈ 3 to 12 km altitude, with accuracy comparable to the
173 sondes (5 to 10%; Leblanc et al., 2018), and slightly coarser altitude resolution (100 m to 2 km).

174 We also use global atmospheric composition re-analyses from the Copernicus
175 Atmosphere Monitoring Service (CAM5, Inness et al., 2019; see also Park et al., 2020), at the
176 grid-points closest to the stations in Table 1. CAM5 re-analyses are based on meteorological

177 fields, and assimilation of satellite observations of ozone and NO₂. They account for the large
178 Arctic stratospheric depletion in spring of 2020 (Manney et al., 2020; Wohltmann et al., 2020),
179 for 2020 meteorological conditions, and for ozone transport, e.g. from the stratosphere to the
180 troposphere (Neu et al., 2014). However, ozone (and NO₂) concentrations in the free troposphere
181 in the CAMS re-analyses are driven primarily by the prescribed emissions. The CAMS re-
182 analyses rely on “business as usual” emissions for 2020, and do not account for COVID 19
183 emission reductions in 2020. Differences between observations (affected by emission reductions)
184 and CAMS re-analyses (“business as usual” emissions) provide a proxy for the effects of
185 COVID 19 emission reductions.

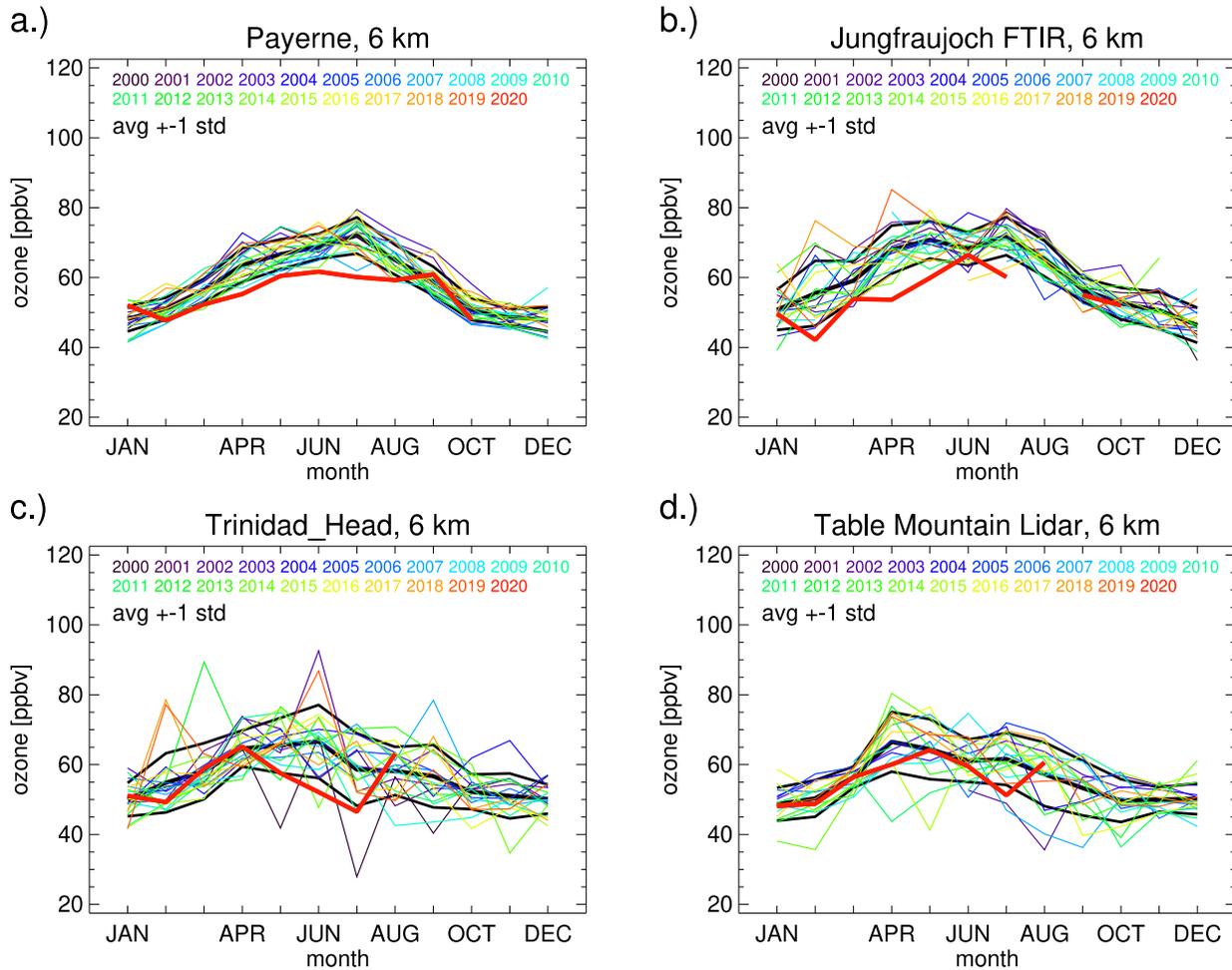
186 (Note: at the time of writing, CAMS re-analyses were available until 12/2019. CAMS
187 operational analyses were used to extend the re-analyses from 01/2020 to 10/2020. For the final
188 version of the paper, CAMS re-analyses will be available until at least 06/2020, and will be
189 used).

190 **3 Results**

191 For selected stations, Fig. 1 presents the annual cycles of tropospheric ozone over the last
192 20 years, at an altitude of 6 km, a representative level for the free troposphere. Monthly means
193 (over 1 km wide layers) reduce synoptic meteorological variability and measurement noise, and
194 focus on longer-term, larger-scale variations.

195 Payerne and Jungfraujoch measure an annual cycle with low ozone in winter and high
196 ozone in summer. This is the case for most stations in the northern extratropics (Cooper et al.,
197 2014; Gaudel et al., 2018; Parrish et al., 2020). Hilo (Hawaii), and Hongkong (both not shown
198 here), further south and in the Pacific region, have an annual cycle where tropospheric ozone
199 peaks in spring. To a lesser degree this is also seen at Table Mountain (California). At tropical
200 stations and in the Southern Hemisphere (not shown), the annual cycle generally peaks in
201 September or October (=spring in the Southern Hemisphere), and has a smaller amplitude
202 (Cooper et al., 2014; Gaudel et al., 2018; Thompson et al., 2012). Increased photochemical
203 production due to more sunlight and warmer temperatures is the main driver for the summer
204 ozone maximum in the northern extratropics (Wu et al., 2007; Archibald et al., 2020).

205 Figure 1 shows substantial variations from year to year. Apart from these variations, Fig.
206 1 shows ozone levels below average in the year 2020 at all four stations (thick red lines in Fig.
207 1). At Payerne and Jungfraujoch, and a number of other stations, monthly means from February
208 2020 through August 2020 were actually the lowest, or close to the lowest, since 2000.
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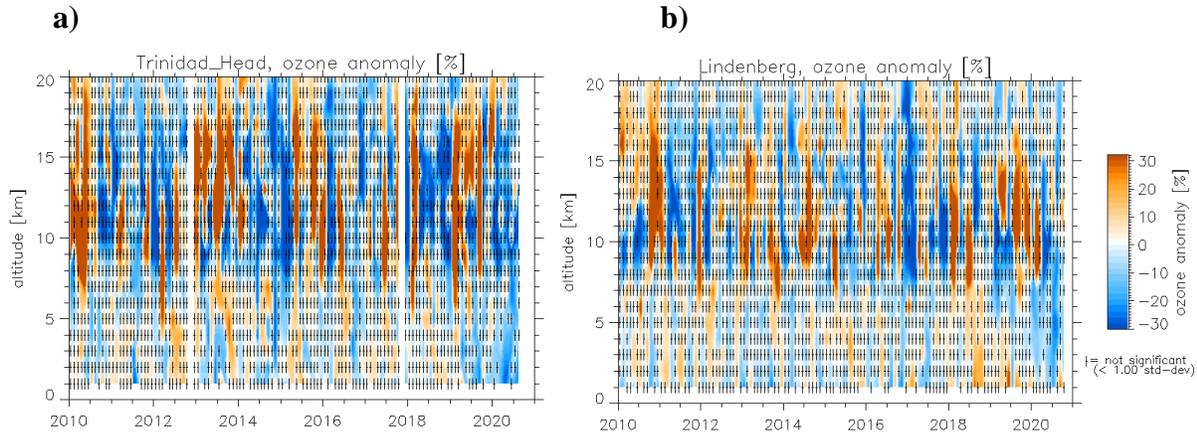
211 **Figure 1.** Observed ozone monthly means, from January 2000 to August 2020, and at four
 212 typical stations. Results are for 6 km altitude. The thick red line highlights the year 2020.
 213 Climatological average, and standard deviation over the years 2000 to 2020 are indicated by the
 214 thick black lines. Payerne (a) and Trinidad Head (c) are sonde stations. Jungfraujoch (b) is an
 215 FTIR station. Table Mountain (d) is a lidar station.

216

217 Ozone anomalies as a function of time and altitude are presented in Fig. 2. For clarity, we
 218 only show the years 2010 to 2020. Both stations in Fig. 2 show varying positive and negative
 219 anomalies at different altitudes. The largest anomalies occur in the 8 to 15 km region, and are
 220 caused by meteorological changes, movement of jet streams, changes in tropopause height and
 221 location, and large variations of the stratospheric circulation (e.g. Neu et al., 2014). In the
 222 troposphere (≈ 1 to 10 km), the largest and most notable negative anomaly at both stations occurs
 223 in 2020 (dark blue region in Fig. 2). This negative anomaly covers several months and most
 224 altitudes from 1 to 10 kilometers. Similar significant, extended negative anomalies throughout
 225 the troposphere occur at many northern extratropical stations in 2020, but are not seen in
 226 previous years, and not across so many locations at the same time.

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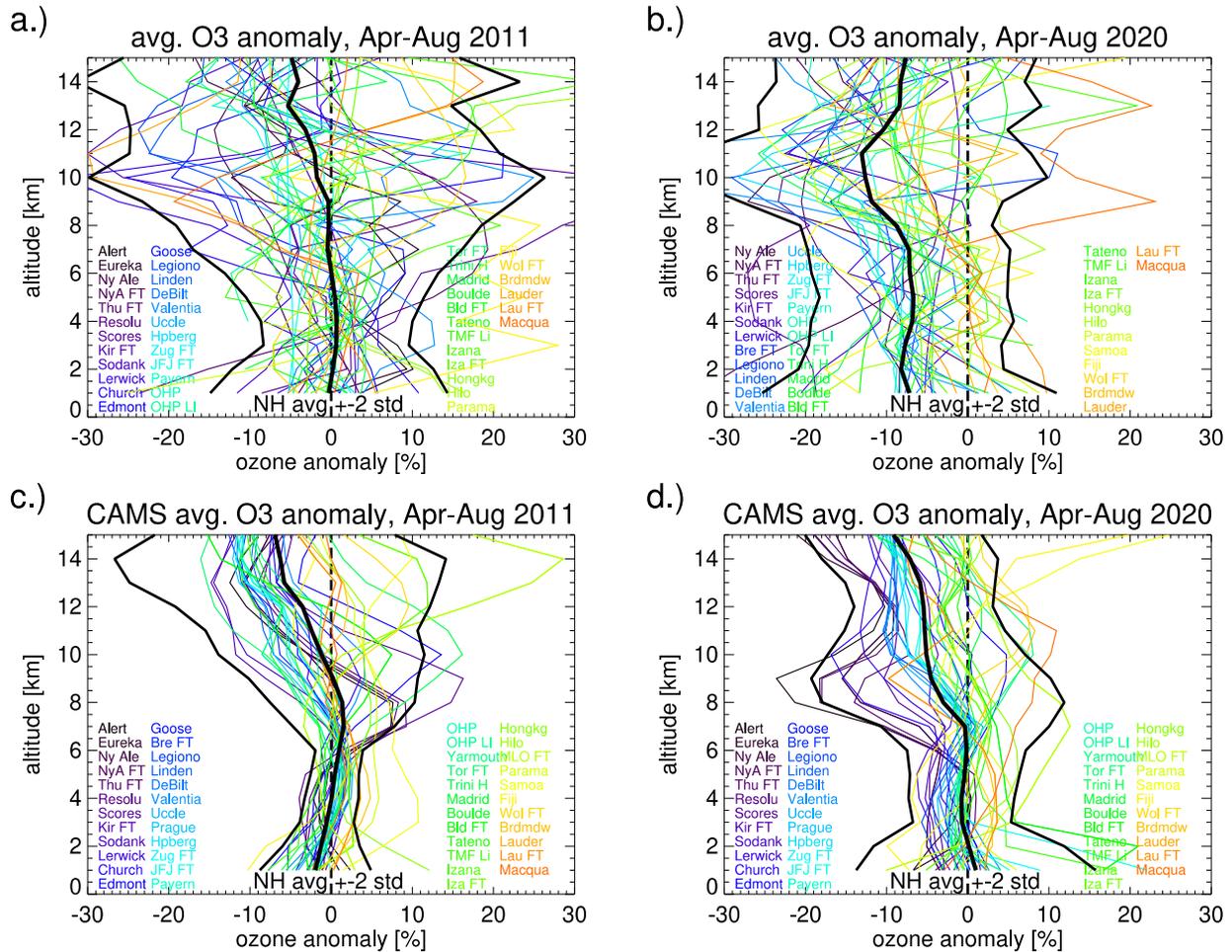
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230 **Figure 2.** Monthly mean ozone anomalies as a function of altitude and time, for the years 2010
 231 to 2020. Years are labeled with tick marks on January 1st. Panel **a)** is for Trinidad Head. Panel **b)**
 232 is for Lindenberg. Anomalies are in percent, relative to the climatological monthly means
 233 calculated for the period 2000 to 2020 (all Januaries, all Februaries, ..., all Decembers).
 234 Anomalies less than 1 standard deviation are crossed out as “insignificant”.

235

236 Figs. 1 and 2 show the largest negative anomalies in the troposphere from April to
 237 August 2020. Therefore, Fig. 3 compares anomaly profiles averaged over those five calendar
 238 months, for the years 2011, and 2020. Both years saw unusually large springtime ozone
 239 depletion in the Arctic stratosphere (Manney et al., 2020; Wohltmann et al., 2020). In the
 240 stratosphere, above ≈ 10 km, the Arctic depletion appears as low ozone, both in observations and
 241 CAMS results, and particularly for the stations north of 50°N . In both stratosphere and
 242 troposphere, the observed profiles are much noisier than the smooth CAMS profiles. In 2020,
 243 most observed single station anomaly profiles (Fig. 3b) are negative throughout the troposphere
 244 (between 1 and 10 km). This is not the case in 2011 (Fig. 3a, 3c). It is also not reproduced by the
 245 CAMS data in 2020 (Fig. 3d).

246 The difference in 2020 is even clearer for the northern extratropical station average
 247 profiles (thick black lines in Fig. 3). The observed 2020 northern extratropical average anomaly
 248 is clearly negative, -6% to -9% from 1 to 8 km (Fig. 3b), throughout much of the troposphere,
 249 whereas in the CAMS data (Fig. 3d) it is close to zero. Fig. 3 indicates that Arctic stratospheric
 250 springtime depletion ozone did not have a large effect on tropospheric ozone in 2011 and 2020,
 251 and that the CAMS “business as usual” simulation does not account for the observed large
 252 negative tropospheric anomaly in 2020.



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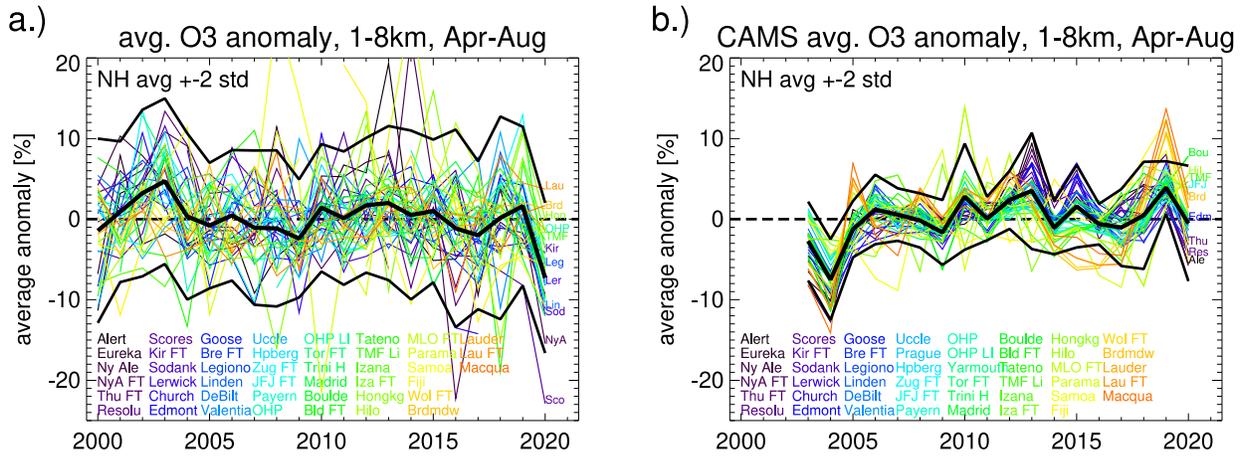
254 **Figure 3.** Ozone anomaly profiles (in percent), averaged over the months April to August.
 255 Stations are excluded in years where their data cover less than three of these five months. Panel
 256 **a)** for the year 2011. Panel **b)** for the year 2020. Colors and stations are sorted by decreasing
 257 latitude. Thick black line: average over all stations north of 15°N (=all stations, except
 258 Paramaribo, Samoa, Fiji, Wollongong, Broadmeadows, Lauder, Macquarie Island). Thin black
 259 lines: ± 2 standard deviations around the average. Panels **c), d)**: Same as a), b), but for
 260 atmospheric composition re-analyses from CAMS at the grid-points closest to the stations.

261

262 Time series of average tropospheric anomalies (averaged from April to August, and now
 263 additionally from 1 to 8 km altitude), are shown in Fig. 4. In the observations (left panel) the
 264 year 2020 stands out with large negative anomalies. This is not seen in the CAMS data. In almost
 265 all twenty previous years, tropospheric ozone anomalies (colored lines) are scattered around
 266 zero. The northern extratropical station average (thick black line) is usually smaller than $\pm 3\%$.
 267 The only other exception is the positive anomaly in the (European) heat-wave summer of 2003
 268 (Vautard et al., 2007) in the observations. (The 2003 and 2004 CAMS results might have a low
 269 bias). The large negative northern extratropical anomaly in the observations in 2020, $\approx -7\%$, is

270 definitely unique in the 21 year observational record, and is not reproduced by the CAMS
 271 “emissions as usual” simulation.

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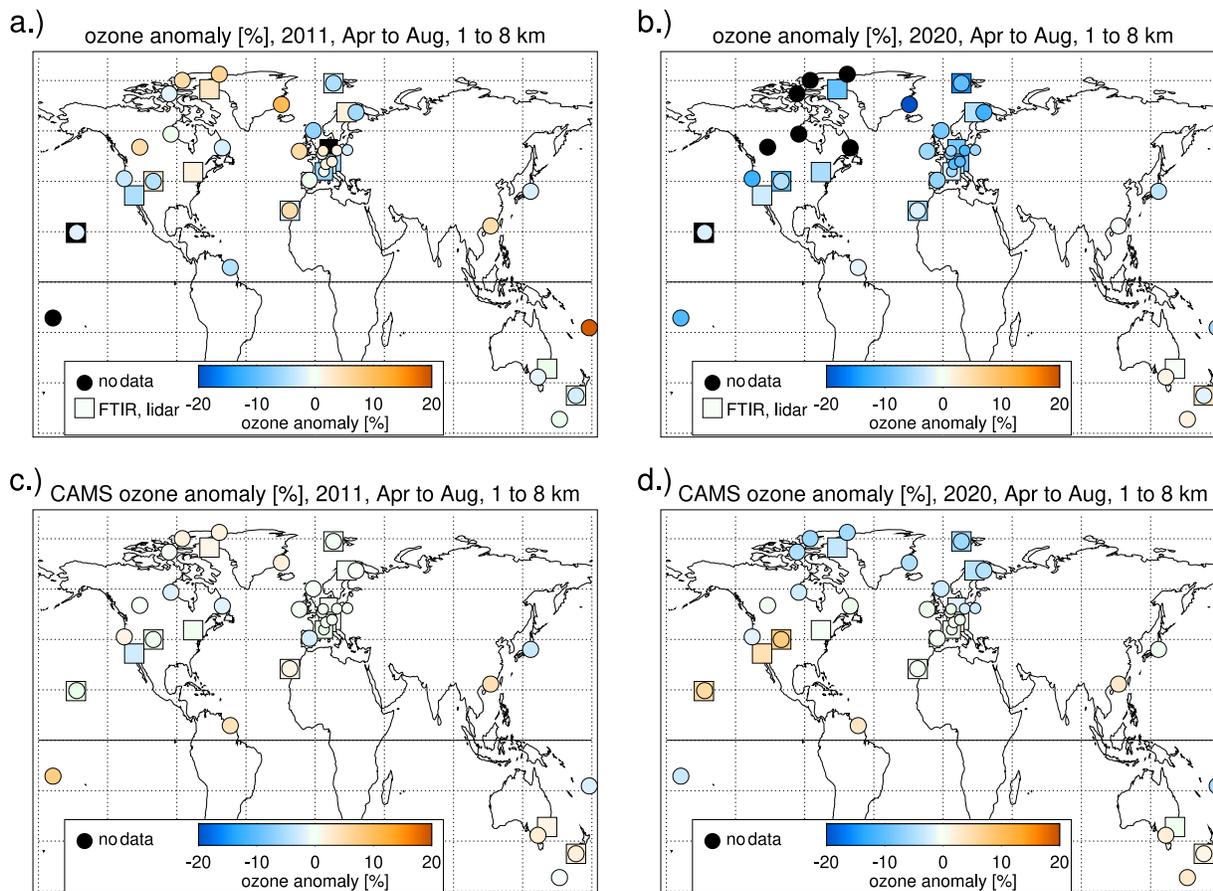


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274 **Figure 4.** Tropospheric ozone anomaly, averaged over the months April to August and over
 275 altitude from 1 to 8 km. Time series for the years 2000 to 2020. Panel a) Results from the
 276 observations. Panel b) same, but for CAMS atmospheric composition re-analyses. Thick black
 277 line: Average over all stations north of 15°N. Thin black lines: ± 2 standard deviations around the
 278 average.

279

280 The geographic distribution of the average tropospheric ozone anomalies is shown for
 281 2011 and 2020 in Fig. 5. 2020 stands out in the observations with large negative anomalies at
 282 nearly all Northern Hemisphere stations, and a fairly uniform geographical distribution (see
 283 Table S1 of the supplement for the numerical values). CAMS does show negative anomalies in
 284 2020, but only north of 50°N, and not as large as the observations. In the Southern Hemisphere
 285 in 2020, agreement between observations and CAMS is quite good. In 2011, some stations show
 286 positive anomalies. Negative anomalies are not as large as in 2020, and the geographical
 287 distribution is less uniform. Agreement between observations and CAMS is reasonable in 2011.



288

289 **Figure 5.** Geographic distribution of observed tropospheric summer ozone anomalies (averaged
 290 over the months April to August, and over altitudes from 1 to 8 km) for the years **a)** 2011 and **b)**
 291 2020. Panels **c)** and **d)**: same, but for CAMS results at the station locations. Colored circles (or
 292 squares) give the anomaly at the ozonesonde stations. Squares are for FTIR and lidar stations.
 293 See Table S1 of the supplement for the numerical values. Black filling indicates insufficient data
 294 in the given year.

295

296 4 Discussion and Conclusions

297 Ozone stations in the northern extratropics indicate exceptionally low ozone in the free
 298 troposphere (1 to 8 km) in spring and summer 2020. Compared to the 2000 to 2020 climatology,
 299 ozone was reduced by 7% (≈ 4 ppbv). Widespread low tropospheric ozone across so many
 300 stations and over several months has not been observed in any previous year since 2000.
 301 Atmospheric composition re-analyses with “business as usual” emissions from the Copernicus
 302 Atmosphere Monitoring Service (CAMS, Inness et al., 2019) do not reproduce the observed low
 303 tropospheric ozone in 2020.

304 The year 2020 stood out in a number of ways: a.) The Arctic stratospheric winter vortex
 305 was exceptionally cold and stable. This produced record levels of springtime ozone depletion in
 306 the Arctic lower stratosphere (Manney et al., 2020; Wohltmann et al., 2020), which might affect

307 tropospheric ozone (Neu et al., 2014). b.) worldwide measures due to the COVID-19 pandemic
308 caused substantial emission reductions in the Northern Hemisphere, up to 60% for some regions
309 and some sectors (Barré et al., 2020; Bauwens et al., 2020; Ding et al., 2020; Guevara et al.,
310 2020). The largest reductions took place in the first months of the year, but air traffic, for
311 example, remains much reduced throughout 2020 (Liu et al., 2020). c.) large wildfires, in early
312 2020 in Australia (Kablick et al., 2020), in August and September 2020 in California, with
313 significant pollution. It is unlikely that the Australian fires have affected tropospheric ozone in
314 the northern extratropics, because pollution from these fires did not reach far into the Northern
315 Hemisphere. The California fires were too late to affect April to July ozone values. In any case,
316 emissions from the wildfires should have increased, not reduced, tropospheric ozone (Archibald
317 et al. 2020).

318 Transport of ozone-depleted air from the Arctic stratospheric vortex appears to be only a
319 minor contributor to the reduced tropospheric ozone: In the observations (Figs. 3 to 5) 2011, and
320 other years with substantial Arctic ozone depletion (2000 and 2016, not shown), do not exhibit
321 large negative anomalies in the troposphere. CAMS atmospheric composition re-analyses also
322 indicate that the 2020 Arctic depletion did not lead to widespread large tropospheric ozone
323 reduction in the northern extratropics (on average less than 1%, see Figs. 3d and 4b). Further
324 evidence for only a small contribution (<1 ppbv, less than one quarter) from the 2020 Arctic
325 depletion to the observed large 7% (or 4 ppbv) reduction comes from the Global Modeling
326 Initiative (GMI) chemistry transport model using MERRA re-analyses (Gelaro et al., 2017;
327 Strahan et al., 2019). See Fig. S2 in the supplement.

328 Weber et al. (2020) recently simulated global effects of COVID-like emission decreases
329 with the UKCA composition climate model. They find tropospheric ozone reductions very
330 similar to our observational results, both in magnitude and in geographical distribution: Figure 2
331 of Weber et al. (2020), for example, shows a fairly uniform ozone decrease by 4 to 7%
332 (depending on emission reduction scenario) in the Northern Hemisphere, and no ozone change in
333 the Southern Hemisphere. This is very similar to our results (e.g., Fig. 5b). Analyses based on the
334 NASA GEOS-CF model also project COVID-19 slowdown-related ozone reductions of about
335 5% for the second half of 2020 (see Fig. 10 of Keller et al., 2020).

336 We suggest that substantial emission reductions caused by COVID-19 pandemic are the
337 major cause for the observed 7% (or 4 ppbv) reduction of northern extratropical free tropospheric
338 ozone in late spring and summer 2020. The large and continuing reduction in air traffic might be
339 particularly important (Grewe et al., 2017).

340

341 **Acknowledgments**

342 The authors greatly acknowledge the know-how and the hard work of station personnel
343 launching the ozonesondes and taking the ground-based measurements. Without their dedicated
344 efforts over many years, and especially during the COVID-19 lockdowns in 2020, investigations
345 like this one are not possible!!

346 **Funding acknowledgments**

347 Deutscher Wetterdienst funds the ozone program at Hohenpeißenberg and makes research like
348 this possible.

349 NOAA GML supported additional launches in Boulder and Trinidad Head in April and May
350 2020. NOAA and NASA's Upper Atmosphere Composition Observations (UACO) Program
351 support the SHADOZ ozone soundings at Hilo, Pago-Pago (American Samoa) and Suva (Fiji).
352 UACO also provides partial support for the Boulder FTIR and the Table Mountain Lidar.

353 The NDACC FTIR stations Bremen, Ny-Ålesund, Izaña, Kiruna, and Zugspitze have been
354 supported by the German Bundesministerium für Wirtschaft und Energie (BMWi) via DLR
355 under grants 50EE1711A, 50EE1711B, and 50EE1711D. Zugspitze has also been supported by
356 the Helmholtz Society via the research program ATMO.

357 The FTIR measurements in Bremen and Ny-Ålesund receive additional support by the Senate of
358 Bremen, the FTIR measurements in Ny-Ålesund also by AWI Bremerhaven. The University of
359 Bremen further acknowledges funding by DFG (German research foundation) TRR 172 – Project
360 Number 268020496 – within the Transregional Collaborative Research Center “Arctic
361 Amplification: Climate Relevant Atmospheric and Surface Processes, and Feedback
362 Mechanisms (AC)3”.

363 The University of Liège contribution has been supported primarily by the Fonds de la Recherche
364 Scientifique - FNRS under grant J.0147.18, as well as by the CAMS project.

365 The Toronto FTIR measurements were supported by Environment and Climate Change Canada,
366 the Natural Sciences and Engineering Research Council of Canada (NSERC), and the NSERC
367 CREATE Training Program in Technologies for Exo-Planetary Science.

368 The University of the Wollongong thanks the Australian Research Council that has provided
369 significant support over the years for the NDACC site at Wollongong, most recently as part of
370 project DP160101598.

371 Supporting results for this manuscript were generated using Copernicus Atmosphere Monitoring
372 Service Information [2000] from the European Community.

373 No author reports a financial (or other) conflict of interest.

374 **Data Sources**

375 Most of the ozonesonde data used in this study are freely available from the World Ozone and
376 UV Data Centre (<https://woudc.org>) at Environment Canada (<https://exp-studies.tor.ec.gc.ca/>),
377 and are downloadable at https://woudc.org/archive/Archive-NewFormat/OzoneSonde_1.0_1/).

378 Some ozonesonde data for 2020 were not yet available at the WOUDC. Instead, rapid delivery
379 data were obtained from <ftp://zardoz.nilu.no/nadir/projects/vintersol/data/o3sondes> (requires
380 registration), at the Nadir database of the Norwegian Institute for Air Quality (NILU,
381 <https://projects.nilu.no/nadir/obs.html>). Registration information, and the same data in a
382 different format, are available from the European Space Agency Validation Data Center
383 (<https://evdc.esa.int/>).

384 For Boulder, Trinidad Head, Hilo, Fiji, and Samoa, stations operated by the US National Oceanic
385 and Atmospheric Administration, Global Monitoring Laboratory
386 (<https://www.esrl.noaa.gov/gmd/ozwv/>), data can be obtained freely from
387 <ftp://aftp.cmdl.noaa.gov/data/ozwv/Ozonesonde/>.

388 FTIR and lidar data, as well as some ozonesonde data, are from the Network for the Detection of
389 Atmospheric Composition Change (<https://ndacc.org>), and are freely available at
390 <ftp://ftp.cpc.ncep.noaa.gov/ndacc/station/> and <ftp://ftp.cpc.ncep.noaa.gov/ndacc/RD/>.
391 Copernicus Atmosphere Monitoring Service (CAMS) global chemical weather re-analyses are
392 available at <https://atmosphere.copernicus.eu/data> . CAMS operational global analyses and
393 forecasts are available at <https://apps.ecmwf.int/datasets/data/cams-nrealtime/> .

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